

Analysis of Flood Routing in Channels

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Abstract - The accuracy of flood routing is an important subject for research in hydrology and hydraulics. Accurate information of the flood peak attenuation and the duration of the high water levels obtained by channel routing are of most importance in flood forecasting operations and flood protection works. This study implements Muskingum method to estimate the inflow and outflow discharge at the river bank. The parameters of Muskingum method are determined using three methods i.e. graphical method, least square method and regression analysis approach. The assessment of results are compared from the above three methods.

Index Terms - Flood routing, hydrological method, Muskingum method, least square method, graphical method, regression analysis.

I. INTRODUCTION

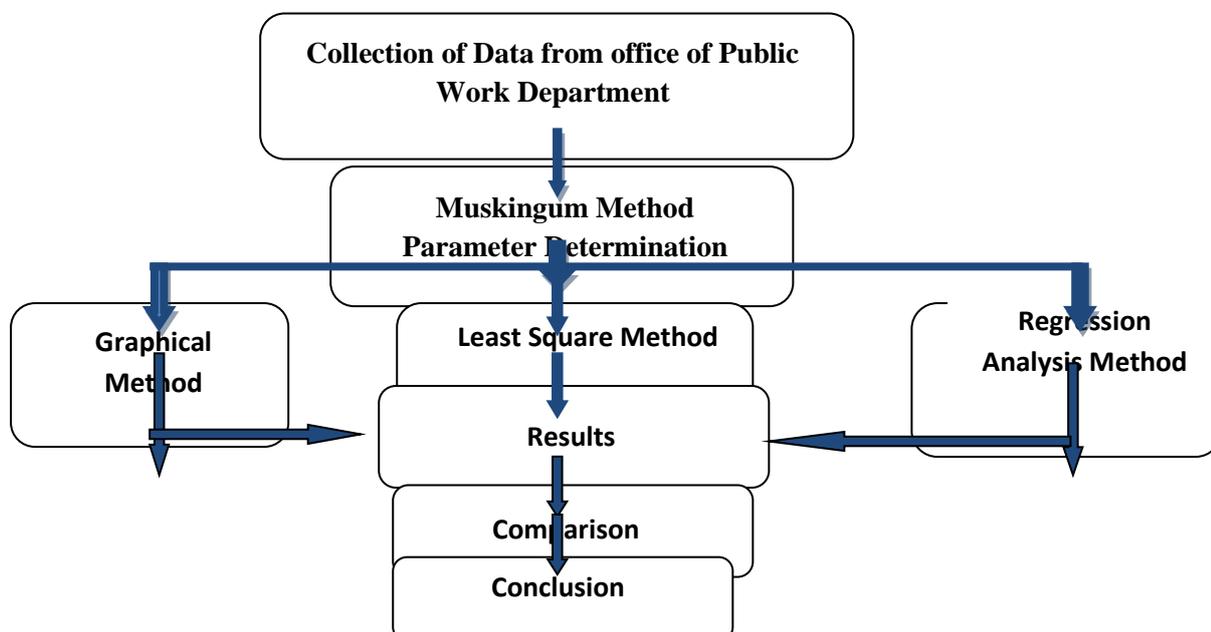
Flood routing is a tool acquiring or creating hydrograph at a specific downstream area of a waterway stream by use of the past information. Flood routing thinks about are valuable in outline of spillway, plan of store, surge determining and surge assurance (Subrmnaya). Surge directing related examinations are used to gauge highest water level streaming in waterway stream. This most extreme water level is profoundly concerned with settling greatest plan release for a spillway. Capacity stores are developed for powerful and dependable instruments of surge controlling measures. This is a flood control technique in which water discharge can be monitored at the downstream end in a given time duration without much effecting the banking areas of a river reach. With a specific end goal to complete above assignments surge directing investigations are required i.e., through reservoir and channel routing. Flood routing is used in (i) flood forecasting, (ii) flood protection, (iii) reservoir design, and (iv) design of spillway and outlet structures. The types of Flood routing are (a) reservoir routing, and (b) channel routing. A variety of routing methods are available and they can be grouped into (1) hydrologic routing, and (2) hydraulic routing. Hydrologic routing methods employ essentially the equation of continuity, on the other hand hydraulic methods use continuity equation along with the equation of motion of unsteady flow (St. Venant equations) hence better than hydrologic methods.

II. OBJECTIVE

To study the newly generating techniques for flood routing and also to evaluate the old used methods. The objectives of this study are to assess the accuracy and reliability of the available flow data collected from the OPWD for modeling, and to check the results through basic flood routing methods.

III. METHODOLOGY USED

This study focuses on the comparison of results calculated from hydrological methods. In this study we have compared results of flood routing inflow and outflow discharges using Muskingum method. Parameters of Muskingum method are determined using three methods namely graphical method, least square method and regression analysis. The reliability of above three methods depends on data availability and the time for calculation. The graphical method is the oldest method to compute the parameters of Muskingum method but the least square method and regression analysis is more fast and accurate in comparison to graphical method as we use mathematical and empirical formulae in least square and regression analysis method.



Hydrologic Channel Routing

The passage of a flood hydrograph through a reservoir or a channel reach is an example of unsteady –flow phenomenon. It is classified in open channel hydraulics as gradually varied unsteady flow phenomenon. The equation of continuity used in all hydrologic routing as the primary equation states that the difference between the inflow and outflow rate is equal ton the rate of change of storage, i.e.

$$I-Q = \frac{ds}{dt}$$

Where I= inflow rate, Q = outflow rate and S = storage.

In channel routing the storage is a function of both outflow and inflow discharges and hence a different routing method is needed. The flow in a river during a flood belongs to the category of gradually varied unsteady flow. The water surface in a channel reach is not only parallel to the channel bottom but also varies with time. Considering a channel reach having a flood flow, the total volume in storage can be considered under two categories as

1. Prism storage
2. Wedge storage

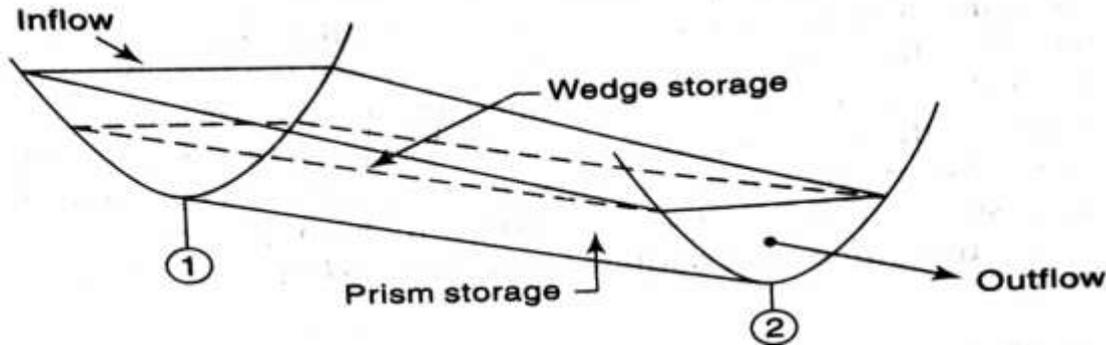


Fig.1. Prism and Wedge Storage In A Channel

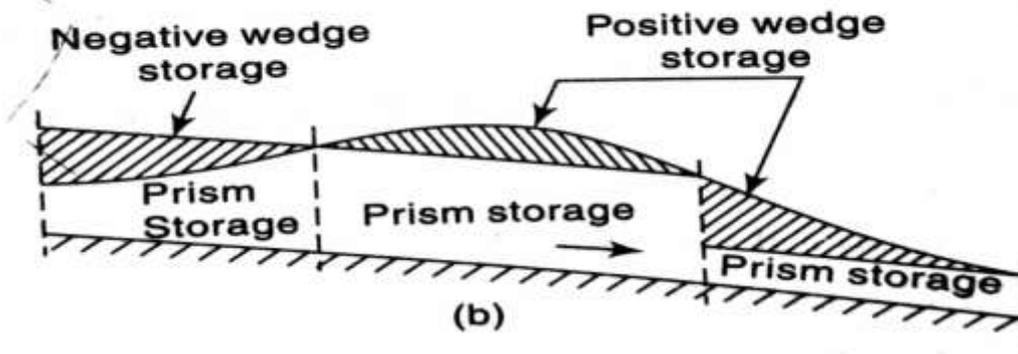


Fig.2. Channel Storage Reach

Hydraulic Methods

Hydraulic flow routing procedures are becoming popular for the purposes of flood routing. This is because hydraulic methods allow flow computation to be varied in both time and space. Hydraulic methods employ the continuity equation together with the equation of motion of unsteady flow (Subramanya, 2009: p280).

The equation of motion for a flood wave is derived from the application of the momentum equation as:

$$\frac{\partial y}{\partial x} + \frac{v}{g} \frac{\partial v}{\partial x} + \frac{1}{g} \frac{\partial v}{\partial t} = s_0 - s_f$$

Where,

v is the velocity of flow at any section in (m/s), S₀ is the channel bed slope and S_f is the slope of the energy line in (m/m).

And the one dimensional continuity equation is given by:

$$\frac{\partial Q}{\partial x} + T \frac{\partial y}{\partial t} = 0$$

The one dimensional continuity and momentum equations mentioned above were first presented by and they are commonly called the Saint Venant equations. If all of the terms in the momentum equation are neglected except for the friction slope (S_f) and bed slope (S₀), the kinematic wave equation simplifies to:

The kinematic wave equation is sufficient for modelling flood waves on steep sloped rivers. When the pressure gradient term is considered, the diffusive wave equation is represented as:

$$S_0 - \frac{dy}{dx} = S_f$$

The shortcomings of these models are represented in the complexity of the methods which are used for solution. These often lead to a numerical instability and problems of 18 convergences. Besides a long running time is needed for the solution which is costly and expensive. Other simplified methods were developed to assist in the calculation, mainly referred to as Hydrological methods.

IV. METHODOLOGIES

Muskingum Method

For a given channel reach by selecting a routing interval Δt and using the Muskingum equation, the change in storage is

$$S_2 - S_1 = K [x (I_2 - I_1) + (1-x) (Q_2 - Q_1)]$$

Where suffixes 1 and 2 refer to the conditions before and after the time interval Δt .

The continuity equation for the reach is

$$(S_2 - S_1) = \left(\frac{I_2 + I_1}{2} \right) \Delta t - \left(\frac{Q_2 + Q_1}{2} \right) \Delta t$$

$$Q_2 = C_0 I_2 + C_1 I_1 + C_2 Q_1$$

$$C_0 = \frac{-Kx + 0.5 \Delta t}{K - Kx + 0.5 \Delta t}$$

$$C_1 = \frac{Kx + 0.5 \Delta t}{K - Kx + 0.5 \Delta t}$$

$$C_2 = \frac{K - Kx - 0.5 \Delta t}{K - Kx + 0.5 \Delta t}$$

C_0 , C_1 , C_2 , and C_3 are dimensionless

A brief review of the Muskingum method precedes the description of the Muskingum-Cunge method. The routing equation is similar for the methods. The Muskingum channel routing method is based on two equations.

The first is the continuity equation or conservation of mass.

$$(S_2 - S_1) = \left(\frac{I_2 + I_1}{2} \right) \Delta t - \left(\frac{Q_2 + Q_1}{2} \right) \Delta t$$

The above Equation states that inflow to the reach minus outflow from the reach is equal to the change in storage. This are the same basic equation used in routing of reservoirs. Where the reservoir routing method assumes relation of storage and outflow discharge, the Muskingum method assumes the amount of storage is related to both inflow and outflow discharge. The Equal reservoir routing method assumes a level pool, and the Muskingum method assumes a sloping water surface.

$$S = K [XI + (1 - X) O]$$

where:

S = reach storage, m^3

K = storage constant, s

X = weighting factor, dimensionless

I = inflow discharge, m^3/s

O = outflow discharge, m^3/s

When X equals zero in the equation reduces to a simple relation of storage and outflow discharge: S is KO (reservoir routing assumption). The values in the equations are in units of meter and seconds. They may also be defined in any units of length and time as long as all values in the equation are consistent.

$$S = K [XI + (1-X) O]$$

$$O_2 = C_0 I_2 + C_1 I_1 + C_2 O_1$$

Coefficients with the sum of C_1 , C_2 , and C_3 equal to 1.0.

An approximation for K is the travel time through the reach or the length of reach divided by the average flow velocity. Either water surface profiles or a solution of Manning's equation are needed for the reach to estimate the average flow velocity. The approximation of K is sensitive to the value of discharge at which velocity is selected. The approximation of K is also sensitive to whether the channel length, floodplain length, or some type of weighted reach length is used. The value of X is between 0.0 and 0.5. A value of 0.0 gives maximum attenuation from the procedure, and 0.5 provides the minimum attenuation.

The routing equation, is applied to a given inflow hydrograph, I_j ($j=1, J$), and initial outflow, O_1 , to calculate the outflow hydrograph, O_j ($j=2, J$), at a downstream section. The constants in the routing equation, C_0 , C_1 , and C_2 are introduced in terms of the channel parameters K and X and the routing time step, t . Limiting the routing time step t within a reasonable range is very important to prevent instabilities in the routing procedure and also to prevent the negative value for coefficient C_1 .

V. Muskingum Method Parameters Determination by Graphical Method:

In the basic Muskingum method, K and X can be graphically estimated from the available inflow and outflow data of the reach of interest. If S is plotted against $XI + (1-X) O$, a straight line with a slope of K should result. Several values of X are tried; the value

that gives the narrowest loop in the plotted relationship is taken as the correct X value and the slope of the plotted relationship is taken as the K value. K is taken as the slope of the straight line of the narrowest loop when $X=0.15$.

The Muskingum method assumes a single stage-discharge relationship. This assumption causes an effect known as hysteresis, which may introduce errors into the storage calculation. The hysteresis effect between reach storage and discharge is due to the different flood wave speeds during the rising and falling limb of the hydrograph. For the same river stage, the flood wave moves faster during the rising limb of the hydrograph.

In spite of its simplicity and its wide applicability, the Muskingum method has the shortcoming of producing a negative initial outflow which is commonly referred to as 'dip' or 'reduced flow' at the beginning of the routed hydrograph. Additionally, the method is restricted to moderate to slow rising hydrographs being routed through mild.

A number of studies have been carried out to analyze flood wave propagation, and the results have shown that the time taken for the center of mass of the flood wave to travel from the upstream end of the reach to the downstream end is equal to K. Thus, K can be easily estimated from the observed inflow and outflow data.

Some factors that are related to a catchment may also play an important role in defining the travel time K.

The surface geology, the soil type, the drainage pattern and the catchment shape may all have influences.

Finally, the Muskingum method also ignores variable backwater effects such as downstream dams, constrictions, bridges and tidal influences.

In small catchments, where measured inflow and outflow hydrographs are not available, or where a significant uncertainty and errors are reported for the outflow data, modeling the flow using this method is quite a source of errors, and the Muskingum method fails to simulate the flow hydrograph using this type of data. In this situation, an alternative procedure developed by Cunge (1967) has received widespread acceptance. This is due to its ability to estimate the model parameters without the observed hydrograph.

In the absence of observed flow data, the Muskingum-Cunge method may be used for parameter estimation. This concept is also applied when measured flow data are available, but with significant degree of uncertainty.

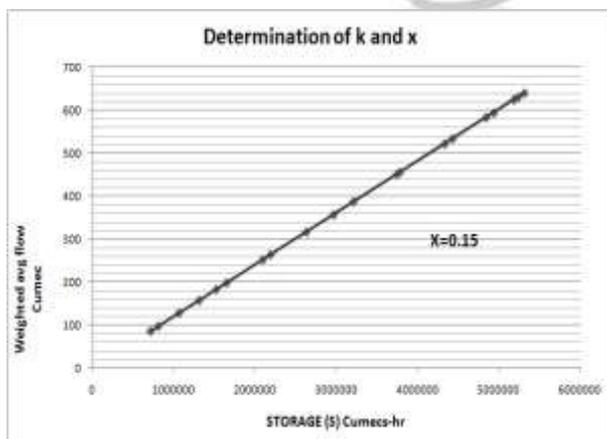
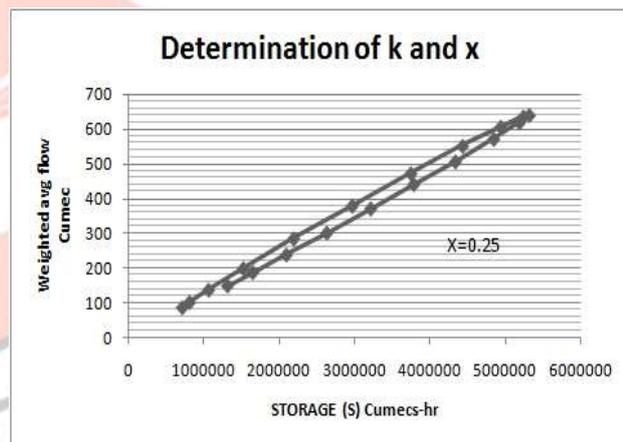
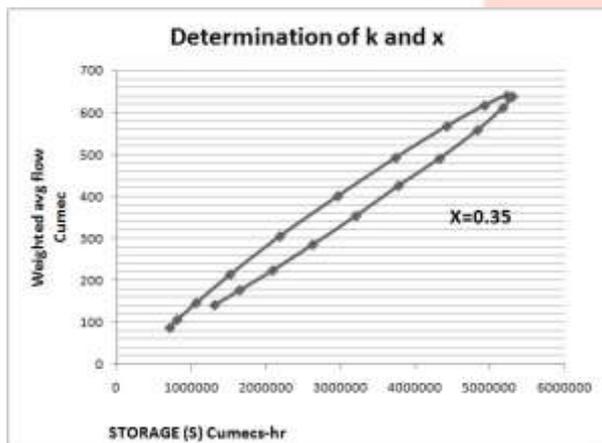
The Muskingum-Cunge parameters are calculated based on the flow and the channel characteristics. This method involves the use of a finite difference scheme for solving the Muskingum equation, where the parameters in the Muskingum equation are determined based on the grid spacing for the finite difference scheme and the channel geometry characteristics.

Now Firstly we can determine parameter using graphical approach for the below data:

1	2	3	4	5	6	7	8	9
					$X=0.25$	$X=0.35$	$X=0.1$	$X=0.15$
<i>Inflow, I</i> (m^3/s)	<i>Outflow, O</i> (m^3/s)	<i>Ave. Inflow</i> (m^3/s)	<i>Ave. Outflow</i> (m^3/s)	<i>Storage</i> (m^3)	<i>Weighted Average Flow</i> $XI + (1-X)O$ (m^3/s)			
93	85			715000	87	87.8	85.8	86.21204
137	91	115	88	812200	102.5	107.1	95.6	97.96922
208	114	172.5	102.5	1064200	137.5	146.9	123.4	128.2414
320	159	264	136.5	1523200	199.25	215.35	175.1	183.3923
442	233	381	196	2189200	285.25	306.15	253.9	264.6645
546	324	494	278.5	2965000	379.5	401.7	346.2	357.634
630	420	588	372	3742600	472.5	493.5	441	451.816
678	509	654	464.5	4424800	551.25	568.15	525.9	534.6043
691	578	684.5	543.5	4932400	606.25	617.55	589.3	595.1201
675	623	683	600.5	5229400	636	641.2	628.2	630.8782

634	642	654.5	632.5	5308600	640	639.2	641.2	640.788
571	635	602.5	638.5	5179000	619	612.6	628.6	625.3037
477	603	524	619	4837000	571.5	558.9	590.4	583.9104
390	546	433.5	574.5	4329400	507	491.4	530.4	522.3653
329	479	359.5	512.5	3778600	441.5	426.5	464	456.2743
247	413	288	446	3209800	371.5	354.9	396.4	387.8502
184	341	215.5	377	2628400	301.75	286.05	325.3	317.2138
134	274	159	307.5	2093800	239	225	260	252.7894
108	215	121	244.5	1649200	188.25	177.55	204.3	198.789
90	170	99	192.5	1312600	150	142	162	157.8796

Now graph between storage (S) in $m^3/s \cdot h$ & $XI+(1-X) O m^3/s$



From the above graph best fit line for least square is for $X=0.15$
Slope of the line for $X=0.15$ is K & given by $K=2.3h$.

VI. Muskingum Method Parameters Determination by Least Square Method:

It consisting in generating graph of $[XI+(1-X) O]$ vs storage (S) for different value of X arbitrarily selected such that $0 < X < 0.5$. the optimal value of X is selected as that which produce narrowest and straightest loop graph of $[XI+(1-X)O]$ vs X. the slope of the least square fit to resulting point is the estimate of K.

$$A = \frac{\sum_{i=1}^n O_i^2 \sum_{i=1}^n S_i I_i - \sum_{i=1}^n O_i I_i \sum_{i=1}^n S_i O_i}{\sum_{i=1}^n I_i^2 \sum_{i=1}^n O_i^2 - [\sum_{i=1}^n O_i I_i]^2}$$

$$B = \frac{\sum_{i=1}^n I_i^2 \sum_{i=1}^n S_i O_i - \sum_{i=1}^n O_i I_i \sum_{i=1}^n S_i I_i}{\sum_{i=1}^n I_i^2 \sum_{i=1}^n O_i^2 - [\sum_{i=1}^n O_i I_i]^2}$$

Now using least square method we find out the value of K & X

Inflow (m ³ /s)	Outflow (m ³ /s)	Storage (S)	O ² (m ³ /s) ²	I ² (m ³ /s) ²	OI (m ³ /s) ²	SO (m ⁶ /s)	SI (m ⁶ /s)
93	85	715000	7225	8649	7905	60775000	66495000
137	91	812200	8281	18769	12467	73910200	111271400
208	114	1064200	12996	43264	23712	121318800	221353600
320	159	1523200	25281	102400	50880	242188800	487424000
442	233	2189200	54289	195364	102986	510083600	967626400
546	324	2965000	104976	298116	176904	960660000	1618890000
630	420	3742600	176400	396900	264600	1571892000	2357838000
678	509	4424800	259081	459684	345102	2252223200	3000014400
691	578	4932400	334084	477481	399398	2850927200	3408288400
675	623	5229400	388129	455625	420525	3257916200	3529845000
634	642	5308600	412164	401956	407028	3408121200	3365652400
571	635	5179000	403225	326041	362585	3288665000	2957209000
477	603	4837000	363609	227529	287631	2916711000	2307249000
390	546	4329400	298116	152100	212940	2363852400	1688466000
329	479	3778600	229441	108241	157591	1809949400	1243159400
247	413	3209800	170569	61009	102011	1325647400	792820600
184	341	2628400	116281	33856	62744	896284400	483625600
134	274	2093800	75076	17956	36716	573701200	280569200
108	215	1649200	46225	11664	23220	354578000	178113600
90	170	1312600	28900	8100	15300	223142000	118134000
			SO ² = 3514348	SI ² = 3804704	SIO = 3472245	SSO = 29062547000	SSI = 29184045000

Example Problem Using Least Square Method:-

Formulas used generally for estimating storage are:

$$S_{i+1} = S_i + \Delta t / 2 (I_{i+1} - I_i) - \Delta t / 2 (O_{i+1} + O_i)$$

Formulas for least Square Approach:

From above we can determine the parameter 'K' as

$$K = A + B \text{ or } X = A / A + B$$

$$A = 1255.626164 \text{ sec. } B = 7079.100513 \text{ sec}$$

$$K = A + B$$

$$K = 8284.726677 / 3600$$

$$K = 2.3 \text{ h}$$

$$X = A / (A + B) = 0.151559154$$

VII. Muskingum Method Parameters Determination By Regression Analysis:

This method is used for rainfall runoff correlation because of several factors effecting runoff resulting from a given rainfall. The relationship between these two is quite complex, in case of flood routing through Muskingum method, the parameters determined are through relation between storage and wetted average flow the coefficient of correlation is found. The coefficient of correlation whose value is minimum, will give the best correlation, however we may use straightlinear regression for small and medium size catchments and exponential form of regression for large catchments.

The general formulae used in linear regression analysis are as follows:

$$R = aP + b$$

Where a and b are constants representing abstractions. The values of a and b are given by the following equations:

$$a = \frac{N(\sum P.R) - (\sum P)(\sum R)}{N(\sum P^2) - (\sum P)^2}$$

$$b = \frac{\sum R - a(\sum P)}{N}$$

$$r = \frac{N(\sum P.R) - (\sum P)(\sum R)}{\sqrt{[N(\sum P^2) - (\sum P)^2] \times [N(\sum R^2) - (\sum R)^2]}}$$

From Data Given Table 1.3

P (storage)	P ² (storage)	X=0.25 R ₁ ²	PR ₁	X=0.35 R ₂ ²	PR ₂	X=0.15 PR ₃	R ₃ ²
715000	5.1122*10 ¹³	7569	62205000	7708.84	62777000	61641608.6	7432.515
812200	6.5966*10 ¹¹	10506.25	83250500	11470.41	86986620	79570600.48	9597.92
1064200	1.132*10 ¹²	18906.25	14632750 0	21579.61	156330980	136474497.9	16445.856
1523200	2.32*10 ¹²	39700.562	30349760 0	46375.622	328021120	279343151.4	33632.7357
2189200	4.7925*10 ¹²	81367.5625	62446930 0	9372.8225	670223580	579403523.4	70047.297
2965000	8.7912*10 ¹²	144020.25	11252175 00	161362.89	1191040500	1060384810	127902.078
37426000	1.4007*10 ¹³	223253.25	17683785 00	243542.25	1846973100	1690966562	204137.6979
4424800	1.9578*10 ¹³	303876.562	24291710 00	322794.42 25	2513950120	236551.7107	285801.7576
4932400	2.432*10 ¹³	367539.0625	29902675 00	381368.00 25	3046003620	2935370381	354167.93
5229400	2.7346*10 ¹³	404496	33258984 00	411137.64	3353091280	329914459	398007.3032

5308600	2.81812×10^{13}	409600	33975040 000	408576.64	3393257120	3401687177	410609.2609
5179000	2.6822×10^{13}	383161	32058010 00	375278.76	3172650400	3238447862	391004.7172
4837000	2.3396×10^{13}	194922.25	21355355 00	213906.25	2237112500	2206998789	208186.23
4329400	1.8743×10^{13}	138012.25	16083721 00	125954.01	1536504060	1679158656	150427.77
37789600	1.4277×10^{13}	91053.0625	11401925 50	8182.6025	1080868530	1198624065	100624.594
3209800	1.0302×10^{13}	57121	76714220 0	50625	722205000	811403416.1	63902.48075
1312600	1.7229×10^{13}	22500	19689000 0	20164	186389200	207232763	24925.9681
$\sum P = 55553000$	$\sum p^2 = 24.2407 \times 10^{13}$	$\sum R_1^2 = 2897607.312$	$\sum PR_1 = 2.532 \times 10^{10}$	$\sum R_2^2 = 297739.6572$	$\sum PR_2 = 2.55843 \times 10^{10}$	$\sum PR_3 = 2.52313 \times 10^{10}$	$\sum R_3^2 = 2856854.111$

Now,

Coefficient of correlation,

for $X=0.25, r_1=0.81177$

For $X=0.35, r_2=0.7985$

For $X=0.15, r_3=0.8329$

Hence, from the above calculation, the best correlation found from the value of $X=0.15, r_3=0.8329$.

VIII. Results and Analysis:

In above data analysis, we have found that all the three methods are equally liable for exact calculation of parameters of Muskingum method. Thereby we can use anyone of the method for the calculation of Muskingum method.

In graphical method, we have found that linear shape of the graph is coming from the value of $x=0.15$ and slope of the line for $x=0.15$ gives value of k as $K=2.3$ hour.

Also in least square method, we have found that the value of X for which the best relationship is found is for $x=0.1515$, value of $K=2.3$ hour.

$K=A+B$ and $X=A/(A+B)$

Where $A=1255.626164$ sec

$B=7079.100513$ sec

From regression analysis, we have found coefficient of correlation for different values of $X=0.25, 0.15, 0.10$

The coefficient of correlation which is nearly equal to 1 found for $x=0.15$, which is $r_3=0.8329$.

Hence, we have found that all the three method gives the same result for $x=0.15$.

Therefore, we can apply any of the above mentioned methods which is suitable condition of given data.

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[REFERENCES]

- [1]. Akan, A. (2006) Open Channel Hydraulics, Oxford: Elsevier Ltd.
- [2]. American Society of Civil Engineers (ed.) (1992) Design and construction of urban irrigation system.
- [3]. Storm-water management systems, New York: American Society of Civil Engineers.
- [4]. A Textbook of Engineering Hydrology by K. Subramanya, TMH Publication.
- [5]. A Textbook of Irrigation and Water Power Engineering by Dr. Pande B B Lal, Laxmi Publication.
- [6]. Kumar, Nagesh, Baliarsingh Falguni, Raju Srinivasa.K, 2010, Extended Muskingum Method for Flood Routing.
- [7]. Carter. W. R, Godfrey. G.R, 1960, Storage and Flood Routing, Manual of Hydrology.
- [8]. Elnbashir. T Safa, 2011, Flood Routing in natural channels using Muskingum methods, Dublin Institute of Technology